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CALCULATION OF ANTHROPOGENIC CARBON IN THE SOUTHERN OCEAN



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Anthropogenic Carbon (Cant)

The sources of anthropogenic carbon are: (1) Fossil fuel emission, (2) Land use change (deforestation, agriculture, cattle farming), and (3) Cement production (as a by-product). The excess of CO_2 due to anthropogenic emissions causes the Earth temperature to rise.

The ocean plays a significant role in absorbing this anthropogenic carbon, and thus mitigating climate change.





Fig. 1: Oceanic carbonate system and ocean acidification process.

Fig. 2: Antarctic continent and Southern Ocean.

TrOCA method

Since it is impossible to directly measure the anthropogenic carbon concentration in the ocean we use the Tracer combining Oxygen, inorganic Carbon and total Alkalinity (TrOCA) method (Touratier et al., 2007). TrOCA is based on (1) a relationship based upon the Redfield's relation (Eq. 1) and (2) and a relation between potential temperature (θ) , oxygen (O₂), total inorganic carbon (C_T) and total alkalinity (A_T).

$$(C_{106}H_{263}O_{110}N_{16}P) + 138O_2 \qquad Eq. 1$$

Respiration,
Decomposition
 \downarrow 106CO₂ + 16NO₃⁻ + HPO₄²⁻ + 122H₂O + 18H⁺

The total inorganic carbon (C_T) varies with depth due to the processes of respiration as well as decomposition of calcium carbonate. It is consumed during photosynthesis in the surface.

Rearranging this equation and using the same concept used by Broecker (1974) when creating the NO and PO tracers, yields the TrOCA equation 2:

$$\begin{bmatrix} Eq. \ 2 \\ TrOCA = \ O_2 + 1.279 \left[C_T - \frac{1}{2} A_T \right] \end{bmatrix}$$

From this equation Cant can be calculated as a function of C_{T} , O_2 , total alkalinity, and potential temperature (θ), as follows (Eq. 3):

$$\begin{bmatrix} \text{Eq. 3} \\ \text{C}_{\text{ant}}^{\text{TrOCA}} = \frac{\text{O}_2 + 1.279 \left[\text{C}_{\text{T}} - \frac{1}{2}\text{A}_{\text{T}}\right] - \text{e}^{\left[7.511 \left(-1.087 \times 10^{-2}\right)\theta - \frac{7.81 \times 10^5}{\text{A}_{\text{T}}^2}\right]}}{1.279}$$

This simple TrOCA method can be applied over large spatial areas since it relies only on physical and chemical properties that are commonly used in oceanography.

This study is part of manuscript which was submitted to the scientic journal Deep-Sea Reseach II and it is currently under review.

Reference

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The Southern Ocean

The Southern Ocean is considered as one of the main sinks of atmospheric CO₂ among all the oceanic basins. Due to low water temperature, strong winds, and seasonal primary production, it absorbs ~40% of the anthropogenic CO₂ from the atmosphere (e.g., Mikaloff Fletcher et al., 2006). The primary production in the Southern Ocean is larger during austral summer when the input of nutrients, necessary for the biological community, is higher due to continent discharges (mostly deglaciation and sea ice melting).

This Study

We present some results of a study which calculated the anthropogenic carbon in a the Gerlache Strait, a costal zone in the Northern Antarctic Peninsula.

Data

- 1996 FRUELA cruise (CDIAC database, Ríos et al. 2007);
- A_T & pH were measured by potentiometric titration and using a pHmeter on NBS scale, respectively;
- C_{T} was estimated from measured A_{T} and pH.
- 2016 NAUTILUS CRUISE (Lencina-Avila et al. under review)
- A_T & C_T were measured by potentiometric titration;

Methodology

- We calculated the anthropogenic carbon (C_{ant}) using the TrOCA method (Touratier et al. 2007) for years 1996 and 2016;
- 2. We subtracted the C_{ant} from the C_T concentrations, yielding the natural concentration of inorganic carbon (C_T°) ;

If the water masses were the same in 1996 and 2016, then we can estimate the penetration of C_{ant} . However, we did observed minor differences in hydrographical conditions, so;

3. We normalized the C_T° and A_T data to the average salinity to correct for carbonate dilution.







ig. 5: Normalized the C₇⁺ and A₇ plot for the respective stations airs between RFUELA and NAUTILUS cruises in (a) the surface rater, (b) the intermediate layer, and (c) the deep layer.



Fig. 4: 9-5 diagrams with inorganic carbon (C,) for (A) 1996 FRUELA and (B) 2016 ANUTILLS, respectively. COW, TBW, and TWW stard for Gircurnpolar Deep Water, Transitional Zonal Water with Bellingshausen Sea influence, and Transitional Zonal Water with Weddell Sea influence, respectively. Water masses indices from Garcia et al. (2002). The dashedthick line marks the mixing zone between South and North Gerlache below MLD.

- There is a considerable increase in C_T concentrations from 1996 to 2016 (Fig. 3e and 3j);
- The southern sector shows higher concentrations of C_T than in the northern sector, which seems related to the hydrography of the Strait;
- The same water masses were identified during both cruises (Fig. 4);
- There is an indication of intrusion of Cant into intermediate (175 m) and deep (700 m) waters (Fig. 5).



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Développement d'un modèle 3D automates cellulaire pour l'approche des phénomènes des feux de forêts

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Modeling and control of the forest fire phenomenon by Cellular Automata Omar Jellouli¹ Abdes Samed Bernoussi¹ Mina Amharref¹

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Motivation



- Main objective: Predict and control complex systems behaviours
- Considered approach: Cellular Automata
- Studied concepts: Spreadability, Vulnerability, Protector Control.
- Particular task: Possibility of protecting a given vulnerable area.

Cellular automata approach

A cellular automaton (CA) $\mathcal{A} = (\mathcal{L}, \mathcal{S}, N, f)$ is given by

- a lattice \mathcal{L} : a regular grid of cells,
- a finite set of states *S*,
- a finite neighborhood **N**(*c*) of size *n*,
- a set of local transition rules which update the state of cell *c* according to the state of its neighborhood *N*(*c*).

 $f: \begin{array}{ccc} \mathcal{S}^{N(c)} & \longrightarrow & \mathcal{S} \\ s_t(N(c)) & \longrightarrow & s_{t+1}(c) \end{array}$

Spreadability and vulnerability

Let ${\mathcal P}$ be a given property specifying the spatial disturbance defined on the CA state.

Consider the set : $\omega_{t_i} = \{c \in \mathcal{L} | \mathcal{P}s_{t_i}(c)\}$

Let σ be a nonempty subset of ${\mathcal L}$ consisting of n_σ cells.

Definition 1
•
$$\mathcal{P}$$
 is spreadable from ω_{t_0} if:
 $\omega_{t_i} \subset \omega_{t_{i+1}}$
• σ is \mathcal{P} -vulnerable if :
 $\exists t_i \in]t_0, t_N[: \sigma \cap \omega_{t_i} \neq \emptyset$

Denote by $\tau_{t_N} = \bigcup_{t_i \in]t_0, t_N[} \omega_{t_i}$ the trajectory of \mathcal{P} . **Proposition 1**

• σ is vulnerable if and only if: $\sigma \cap \tau_{t_N} \neq \emptyset$

Protector control

- Autonomous system : $\mathcal{A} = (\mathcal{L}, N, \mathcal{S}, f)$
- Disturbed system : A_p=(L, N, S_p, f_p)
 Disturbed-Controlled system
- Disturbed-Controlled system $\mathcal{A}_{pu} = ((\mathcal{L}, \mathcal{S}_{pu}, N, f_p), u)$

 $\begin{array}{l} \sigma \ \subset \ \mathcal{L} \ \text{is vulnerable} \ \leadsto \ \forall c \ \in \ \sigma, \exists t_s \ \in \\]t_0, t_N[, \forall t_s > t_i : s_{t_s}(c) \neq s_{t_s}^{p,0}(c) \end{array}$

Considered control: Let u, applied in $\mathcal{L}_u \subset \mathcal{L}$ (control set) to protect σ :

Denote by $\tau^u_{t_N}$ the controled trajectory of \mathcal{P} . **Definition 2**

- σ is exactly protectable during $]t_0, t_N[$ if : $\exists u \mid \forall t_s > t_i, \forall c \in \sigma , s_{t_i}(c) = s_{t_i}^{p,u}(c)$ (4)
- σ is weakly protectable if:

$$\exists u \mid \forall \varepsilon > 0 : \frac{Card\left(\sigma \cap \tau_{t_N}^u\right)}{Card\left(\sigma\right)} \le \varepsilon \quad (5)$$

Proposition 2

 $s_{t_i}^p$

• σ is protectable, if: $\exists u \mid \tau_{t_N}^u \cap \sigma = \emptyset$ (6)

Control protector approch

Approche

- Stop the disturbance spread, we ensure the desired state, which is manifested in the elimination of the perturbation.
- Change the disturbance trajectory which targets vulnerable areasin order to protect it.

Problem 1 Find $\tau_{T_N}^{p,u}$ with aim : $\tau_{T_N}^{p,u} \cap \sigma = \phi$

considered control:

- Passive control: impliment an action support L_u ⊂ L of the control before the disturbance launch, by knowing the vulnerable zones σ to be proteced, and by anticipating appropriate measures in advance.
- Dynamic control : the implementation of action support is performed during the disturbance spread.

Simulation results

Passive control

- Barrier firebreak intended to slow or block the fire,

- Road networks facilitating the movement of firefighters, emergency personnel maintenance or monitoring, etc.

Dynamic control

(1)

(2)

(3)

Firefighters Shifting,
Reinforcement by the FRA of the Canadian aircraft fleet "Amphibious" aircraft with a capacity



Figure 2: Effective Dynamic Control

Case study : Forest fire dynamics

Disturbed system: CA with Fire
The states set
$$S_p = \{0, 1, 2, 3, 4\}$$
:

 $\mathbf{a}_{or} \mathbf{P}_{or} \mathbf{P}$ State 0 : No fire (just vegetation)

- State 1 : Excited with fire
- State 2 : Fire
- State 3 : Ash
- State 4 : Empty cell (nothing at all)

The transition rules f_p :



Controlled CA

Control effect $u \rightsquigarrow \forall c_{ij} \in \mathcal{L}_u$, $\forall t_i \in I$, $u_{t_i}(c_{ij}) = u$, where u is the performed action to stop the fire spread toward the region σ .





Figure 1: Passive control using firewall trenches



Figure 3: Inefficient dynamic control



Figure 4: Effective Dynamic Protector Control

The simulations illustrate the possibility of controling the fire in Fig. 1 Fig. 2 otherwise (Fig. 3 case) protecting a given region against the fire spread with a 'good' choice of the domain \mathcal{L}_u in Fig. 3.

Work in progress

Theoretical questions :

- Characterize the optimal protector control,
- The observation and the control implementation delay,
- The stochastic CA.

- **Practical questions :**
 - Control implementation in a real case using satellite Images [2],
 - Improve the developped simulator,
 - 3D visualisation.

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Modeling of viscoelastic waves with fractional derivatives

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APPLICATIONS



OBJECTIVES

Powerful numerical methods: **Realistic, Fast, Precise and Robust**

MATHEMATICAL ANALYSIS

Model problem

We consider a class of generalized fractional Zener model in d dimension for waves propagation in dissipative media. Our goal is to determine the displacement u and the stress tensor σ which verify :

$$\begin{cases} \rho(x)\frac{\partial^{2}\boldsymbol{u}}{\partial t^{2}}(x,t) - div\,\sigma(x,t) = f, & \text{in } \mathbb{R}^{d} \times]0, T[, \\ \sigma(x,t) + \tau_{0}(x)\frac{\partial^{\alpha}\sigma}{\partial t^{\alpha}}(x,t) = C(x)\varepsilon(\boldsymbol{u}(x,t)) + \tau_{0}(x)D(x)\varepsilon(\frac{\partial^{\alpha}\boldsymbol{u}}{\partial t^{\alpha}}(x,t)), & \text{in } \mathbb{R}^{d} \times]0, T[, \\ \boldsymbol{u}(x,0) = \boldsymbol{u}_{0}, \frac{\partial \boldsymbol{u}}{\partial t}(x,0) = \boldsymbol{u}_{1}, \sigma(x,0) = \sigma_{0}, & \text{in } \mathbb{R}^{d}. \end{cases}$$

$$(1)$$

Where div is the divergence operator, ε the strain tensor, f is the source density, ρ are physical parameters and τ_0 is relaxation times, C and D two tensors of order 4 symmetric, positive definite

• $D_t^{\alpha} = \frac{\sigma^2}{\partial t^{\alpha}}$ is the fractional derivative of order α :

$$D_t^{\alpha}g(t) = \frac{1}{\Gamma(1-\alpha)} \int_0^t \frac{1}{(t-\tau)^{\alpha}} \frac{\partial g}{\partial \tau}(\tau) d\tau, \quad \alpha \in]0,1[,$$
mma, defined by :

where Γ is the function ga

$$\Gamma(x) = \int_{0}^{+\infty} t^{x-1}e^{-t} dt.$$

Existence and uniqueness of the strong solution

To show the existence and uniqueness of the strong solution, we use the Hille-Yosida theory. To use this theorem, we must transform the model problem (1) in the form of a first order evolution system, for this reason we introduce the following auxiliary differential equation:

$$\begin{cases} \frac{\partial \varphi}{\partial t}(x,t,\xi) = -\xi \varphi(x,t,\xi) + \mathbf{s}(x,t), \text{ in } \mathbb{R}^d \times]0, T[\times [0,+\infty[\,,\\\varphi(x,0,\xi) = 0, & \text{ in } \mathbb{R}^d \times [0,+\infty[\,.\end{cases}) \end{cases}$$

We consider the functional spaces:

$$\begin{split} & \left| \begin{array}{l} \mathcal{L}^{sym}(\mathbb{R}^d) = \left\{ \sigma \in \mathcal{L}(\mathbb{R}^d) / \sigma_{ij} = \sigma_{ji} \forall i, j = 1, ..., d \right\}, \\ & L^2(\mathbb{R}^d, \mathcal{L}^{sym}(\mathbb{R}^d)) = \left\{ \sigma : \mathbb{R}^d \mapsto \mathcal{L}^{sym}(\mathbb{R}^d) / \int_{\mathbb{R}^d} |\sigma|^2 \mathrm{d}x < \infty \right\}, \\ & H^{sym}(div; \mathbb{R}^d) = \left\{ \sigma \in L^2(\mathbb{R}^d, \mathcal{L}^{sym}(\mathbb{R}^d)) / div\sigma \in [L^2(\mathbb{R}^d)]^d \right\}, \\ & \mathbb{H}^{sym}_\alpha = L^2(\mathbb{R}_+, dM_\alpha(\xi)), \\ & \bar{V}^{sym}_\alpha = L^2(\mathbb{R}_+, \xi dM_\alpha(\xi)). \end{split}$$

With $dM_{\alpha}(\xi) = \frac{\sin(\alpha \pi)}{\xi} \xi^{-\alpha} d\xi$. Theorem

For all initial conditions $(u_0, u_1, \sigma_0) \in [H^2(\mathbb{R}^d)]^d \times [H^1(\mathbb{R}^d)]^d \times H^{sym}(div; \mathbb{R}^d)$ and all $f \in C^1(0, T; [L^2(\mathbb{R}^d)]^d)$, there exists a unique solution (u, σ) of the problem (1) which satisfies:

 ${}^{\bullet} {\boldsymbol{u}} \in C^0(0,T; [H^2(\mathbb{R}^d)]^d) \cap C^1(0,T; [H^1(\mathbb{R}^d)]^d) \cap C^2(0,T; [L^2(\mathbb{R}^d)]^d),$ $\sigma \in C^0(0, T; H^{sym}(div; \mathbb{R}^d)) \cap C^1(0, T; L^2(\mathbb{R}^d, \mathcal{L}^{sym}(\mathbb{R}^d))).$

Energy decay result

In order to show that our model is dissipative, we point decreasing quantity that will be called energy of the model. Let (u, σ) be the strong solution of the problem (1). We define the energy of (u, σ) at time t: 1...

$$\boldsymbol{E}(t) = \frac{1}{2} \left(\left\| \frac{\partial \boldsymbol{u}}{\partial t} \right\|_{\rho}^{2} + \|\varepsilon(\boldsymbol{u})\|_{C}^{2} + \|\boldsymbol{s}\|_{Z^{-1}}^{2} + \int_{\mathbb{R}^{d}} \tilde{Z}^{-1} \|\varphi\|_{\tilde{V}_{\alpha}^{sym}}^{2} \right)$$

• $s = \sigma - C\varepsilon(u)$. • Z = D - C is positive definite.

• $\tilde{Z} = Z \tau_0$.





Theorem

The energy E(t) associated to the model problem (1) satisfies this identity: $\left(f, \frac{\partial u}{\partial t}\right) - \int_{\mathbb{T}^d} \tilde{Z}^{-1} \parallel \frac{\partial \varphi}{\partial t} \parallel^2_{H^{sym}_\alpha} dx.$

The energy decreases in the absence of a source term (f = 0).

NUMERICAL ANALYSIS

In this part, we present a numerical analysis of homogeneous problem in 1D cases

$$\begin{cases} \rho \frac{\partial^2 u}{\partial t}(x,t) - \frac{\partial \sigma}{\partial x}(x,t) = f(x,t), & \text{in }]0,1[\times]0,T[,\\ \tau_0 \frac{\partial s}{\partial t}(x,t) + \int_0^{+\infty} (s-\xi\varphi) \, dM_\alpha(\xi) = \mu(\tau_1 - \tau_0) \frac{\partial^2 u}{\partial x \partial t}, & \text{in }]0,1[\times]0,T[\times[0,+\infty[,\\ \frac{\partial \varphi}{\partial t}(x,t,\xi) = -\xi\varphi(x,t,\xi) + s(x,t,\xi), & \text{in }]0,1[\times]0,T[\times[0,+\infty[,\\ \sigma(x,t) = s + \mu \frac{\partial u}{\partial x}, & \text{in }]0,1[\times]0,T[,\\ u(x,0) = u_0, \frac{\partial u}{\partial t} = u_1, \sigma(x,0) = \sigma_0, \varphi(x,0,\xi) = 0, & \text{in }]0,1[\times[0,+\infty[,\\ u(0,t) = u(1,t) = 0, & \text{in }]0,T[. \end{cases}$$

Where μ and ρ are physical parameters and τ_0 and τ_1 are relaxation times For the numerical approximation of the model problem, we use a totally explicit centered finite difference scheme of order 2

$$\begin{cases} u_{i}^{n+1} = \frac{\Delta t^{2}}{\rho} \left(f_{i}^{n} + \mu \frac{u_{i+1}^{n} - 2u_{i}^{n} + u_{i-1}^{n}}{h^{2}} + \frac{s_{i+\frac{1}{2}}^{n} - s_{i-\frac{1}{2}}^{n}}{h} \right) + 2u_{i}^{n} - u_{i}^{n-1}, \\ s_{i+\frac{1}{2}}^{n+1} = \frac{\Delta t}{\tau_{0} + \lambda\Delta t} \left(\frac{\mu(\tau_{1} - \tau_{0})}{h} \left[\frac{u_{i+1}^{n+1} - u_{i+1}^{n}}{\Delta t} - \frac{u_{i}^{n+1} - u_{i}^{n}}{\Delta t} \right] + \sum_{j=1}^{N_{s'}} \tilde{w}_{j} \varphi_{i+\frac{1}{2},j}^{n} + \right) - \frac{\lambda\Delta t - \tau_{0}}{\lambda\Delta t + \tau_{0}} s_{i+\frac{1}{2},j}^{n}, \\ \varphi_{i+\frac{1}{2},j}^{n+1} = \frac{2 - \xi_{j} \Delta t}{2 + \xi_{j} \Delta t} \varphi_{i+\frac{1}{2},j}^{n} + \frac{\Delta t}{2 + \xi_{j} \Delta t} (s_{i+\frac{1}{2}}^{n+1} + s_{i+\frac{1}{2}}^{n}), \end{cases}$$
with $w_{j} = \frac{\sin(\alpha\pi)}{\pi} \xi_{j}^{-\alpha} \Delta\xi, \forall j = 2, ..., N_{s} - 1. \ \tilde{w}_{j} = \xi_{j} w_{j} \frac{2}{2 + \xi_{j} \Delta t} \text{ and } \lambda = \sum_{j=1}^{N_{\xi}} \frac{w_{j}}{2 + \xi_{j} \Delta t}. \end{cases}$

Theorem

The numerical scheme (2) is
$$L^2$$
 stable if, and only if,
 $\Delta t \leq h \left(c\sqrt{\frac{\tau_1}{\tau_0}}\right)^{-1}, \ c = \sqrt{\mu/\rho}.$

NUMERICAL RESULTS

The domain of computation is the unit segment]0, 1[. The exact solution (u, σ) of the problem is computed by the method of separation of variables.



On the left: The numerical solution and the exact solution for $\alpha = 1/2$

CONCLUSION AND PERSPECTIVES

- Realistic model with fractional derivative
- Energy decay and existence and uniqueness of strong solution
- A stable explicit scheme
- The study is extended to 2D and 3D cases.

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Modélisation de l'érosion hydrique des sols par automates cellulaires

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Figure 1: Phénomènes rares sont fréquents

- Suivre l'évolution de l'érosion.
- Prédir et protéger les zones vulnérables.
- Cadre: théorie des systèmes.
- Approche: automates cellulaires.



Figure 2: Manifestation physique de l'érosion

- Erosion: détachement, transport, dépôt.
- Variation: fonction du cycle de l'eau.
 - perte en sol: précipitation,
 - gain en sol: évaporation.
- Facteurs: milieu physique et climat.



Conclusion et perspectives

- Description des processus de l'érosion.
- Construction d'un modèle 2D du suivi de l'érosion.
- Illustration de quelques simulations du modèle.
- Simulation concrètre dans une région réelle
- Calage et validation du modèle.
- Couplage avec les changements climatiques.

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Résultats Les entrées



Signification des états $\{0, 1, 2, 3\} \times \{-1, 0, 1\}$

- \equiv non saturation et pas d'eau surfacique
- saturation et pas d'eau surfacique, 1 \equiv
- 2 non saturation et eau surfacique, \equiv
 - saturation et eau surfacique, =
 - -1 sol érodé, \equiv 0
 - sol initial. \equiv 1
 - ≡ sol déposé.



Modélisation de l'impact des changements climatiques sur les ressources en eaux par Automates Cellulaires : Etude comparative sur deux périmètres au Maroc (Gharb et Loukkous)

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NTRODUCTION Le climat constitue pour le système terrestre un caractère intrinsèque **CONTEXTE DE L'ÉTUDE Problématique** Objectifs généraux de son évolution. Cependant, au cours des dernières décennies il subit de nombreuses variations significatives par rapport aux variations Assurer une gestion durable des ressources en eau et subvenir à la demande croissante en eau destinée à l'irrigation. Comment et à quel point, les changements climatiques affecteront ils la répartition temporelle des ressources en eau dans le futur ? antérieures. Changement climatious Est-ce que ce changement aura le même effet sur les deux périmètres d'étude? Tester le modèle 2CAFDYM sur d'autres zones d'étude Diminution des précipitatio Présentation des deux zones d'étude Cycle hydrologi Superficie : 2.560 km² ÷ Evolution du cycle hydrologique et Type du climat : C Evolution du climat Pmoy : 725 mn Tmoy :11°C en -311 200002 A Sols : 4 types (ss, sa, sH, sPE 10000 Modélisation des Superficie : 6.160 km impacts Type du climat : Climat méd 40000 Quel modèle utilisé? <u>Pmoy :</u> 480 mm. <u>Tmoy :</u>13°C en Hiver et 27°C en Été 0000 Topographie : Relativement plate Sols : 8 types (ss, ss, ss, ss, ss, sre, ms, sc, ss, e de la re Figure : Localisation géographique des deux zones d'étude et emplacement souterraines des stations météorologiques MÉTHODOLOGIE Approche automate cellulaire Approche méthodologique Modèle spatial dynamique, inventé par John von Neumann dans les Ouputs PHASE 1 années 50 Raster MNT Découpage Fusion des images Carte Mod Image fusion (MNT) Automate Cellulaire ique de Terrain Périmètre .shp Digitalisation des Digitalisation de la Georéférencement Acquisition de la carte de la zone Discrétisation (spatiale : espace décomposé en une grille de d'étude types de sol limite du périmètre de la carte cellules ; temporelle: évolution par pas de temps discrets). opriétés de Parallélisme (les cellules évoluent simultanément et de manière interactive) Classification supervisée Reprojection et Pre-proccessing et de la végétation (Rando forest) correction atmosphérique Localité (chaque cellule évolue en fonction de son propre état et de celui d'un ensemble fini de cellules voisines). découpage des (Landsat 4-5, Landsat 7, Landsat 7, Landsat 4-5, Landsat 7, Landsa Landsat 8) Carte d pation du sol images Adaptation (II s'adapte parfaitement à l'architecture des Inputs ordinateurs et aux réalités des systèmes physiques) Modèle 2CAFDYM Outputs Zone Le modèle Two Scale Cellular Automaton for Flow DYnamics Modelling MNT: Modèle Eau évaporée PHASE 3 Gharb PHASE 4 (2CAFDYM), développé par [1], a été utilisé pour modéliser et évaluer l'évolution du cycle de l'eau. Numérique de Terrain Eau ruisselée Type du sol/RFU_{max}. Type d'occupation des sols Eau infiltrée 2CAFDYM(L,N,S,f) Outputs Zone Données climatiques de référence. Condit $f: S^m \rightarrow S \text{ et } S^{t_r} \rightarrow S^{t_{r_r}}$ Ressources en eau superficielle Loukkos $L = \{c_{ij}; i, j \in \mathbb{Z}\}$ (2) limites Ressources en eau souterraine Données climatiques projetées selon les scénarios de [2] $N(c_{ij}) = \{c_{kl} \in L; d(c_{ij}, c_{kl}) \le 1\}$ (3) $S = \{0, 1, 2, 3\}$ (4) Ouputs PHASE 2 Inputs L : représente le treillis. De tel sorte que pour chaque cellule c_{ij}, nous avons toutes les informations nécessaires sur les PPT et les PC. Cette cellule c_{ij} est définie par les Traitement des Moyennes mensuelles, Analyse de la Calcul des paramètr Acquisition des données climatiques . coordonnées (i,j) et une forme carrée ayant une taille de 30x30m. saisonnières et annuelles variabilité climatique données + N : Le voisinage d'une cellule c_{ij} qui affecte l'évolution du c_{ij} dans le temps. Il est de Pas de temps: décadaire Tendances mensuelles Analyse des Détection rayon de 1 et de taille m{3,4,5,6} selon les conditions limites. Avec d est la distance saisonnières et annuelles tendances des ruptures Cartes des tendances dans L x L équivalente à L_{∞} Tendances projetées saisonnières et annuelles annuelles et saisonnières pou Obtent on des tendance Superposition a S : L'état de la cellule, définie comme suit Georéférenc la période 2021-2050 [2] pour chaque station la zone d'étude de la carte $\begin{array}{l} 0 \ , non saturation \left(Gw_{ij}^{\left \lfloor t_{1} \right \rfloor} < S_{ij} \right) et pas d'eau de surface \left(Sw_{ij}^{\left \lfloor t_{1} \right \rfloor} = 0 \right) \\ 1 \ , saturation \left(Gw_{ij}^{\left \lfloor t_{1} \right \rfloor} = S_{ij} \right) \ et pas d'eau de surface \left(Sw_{ij}^{\left \lfloor t_{1} \right \rfloor} = 0 \right) \\ 2 \ , non saturation \left(Gw_{ij}^{\left \lfloor t_{1} \right \rfloor} < S_{ij} \right) et \ avec eau de surface \left(Sw_{ij}^{\left \lfloor t_{1} \right \rfloor} > 0 \right) \\ 3 \ , saturation \left(Gw_{ij}^{\left \lfloor t_{1} \right \rfloor} = S_{ij} \right) \ et \ avec eau de surface \left(Sw_{ij}^{\left \lfloor t_{1} \right \rfloor} > 0 \right) \\ \end{array}$ Matériel *S* = NetBeanside 8.0.2 Délimitation des deux zones d'étude Traitement des données météorologiques. Graphes de la répartition temporelle des précipitations et température moyenne de l'air (périmètre du Gharb). Préparation des cartes d'élévation (MNT) pour les deux zones d'étude Elaboration des cartes de type de sol , de RFU_{max} et d'occupation du sol Cartes de la répartition spatiale de la tendance des précipitations/ température moyenne de l'air annuelles et saisonnières. pour le périmètre du Gharb Redimensionnement des rasters (.tif) et leur conversion en fichier (.txt). PERSPECTIVES FÉRENCES [1] Kassogué, H., Bernoussi, A., Amharref, M., & Ouardouz, M. (2016, September). Modelling of Climate Change Impact on Water Resource Cellular Automata Approach. In International Conference on Cellular Automata (pp. 280-290). Springer International Publishing Comparer les résultat des deux périmètres. [2] Driouech, F. (2010). Distribution des précipitations hivernales sur le Maroc dans le cadre d'un changement climatique : desce et incertitudes. Thèse de l'Université de Toulouse E

Resolution by natural sub-domains of problems with multiphysical contact

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INTRODUCTION

Solving industrial problems using the finite element method usually leads to the assembly of large systems. This one is all the more important in a context of multiphysical problems that involve phenomena complexes of different nature and strongly coupled. For example, when it comes to the study of the thermo-electromechanical behavior of an electrolysis cell which comprises a multitude of contact interfaces of various behaviors. To solve this kind of problem, we usually use methods known as domain decomposition which consists of sharing the domain of computation in sub-domains and share the calculation task.

OBJECTIVES

Complete gear box (provided by Pratt and Whitney Canada).



Great items (gure provided by Pratt and Whitney Canada).



MATHEMATICAL ANALYSIS Conventional definition of the contact interface



Figure 1: geometry of the contact interface

New definition of the contact interface



For the figure, it is easy to show that the corresponding linear system can be written in the form:

Figure 2: Geometry of new contact interface

$$\begin{pmatrix} K_{ie}^{A} & K_{ie}^{A} & 0 \\ K_{ei}^{A} & K_{em}^{A} & K_{em}^{A} \\ 0 & K_{me}^{A} & K_{mm}^{A} \end{pmatrix} = \begin{pmatrix} \Delta U_{i}^{A} \\ \Delta U_{e}^{A} \\ \Delta U_{m}^{A} \\ \Delta U_{m}^{B} \\ 0 \\ 0 \\ 0 \\ K_{em}^{B} & K_{eB}^{B} \\ K_{ie}^{B} \\ K_{ie}^{B} \\ K_{ie}^{B} \\ \Delta U_{b}^{B} \\ \Delta U_{b}^{B} \end{pmatrix} = \begin{pmatrix} R_{i}^{A} \\ R_{e}^{A} \\ R_{m}^{A} \\ R_{m}^{B} \\ R_{e}^{B} \\ R_{e}^{B} \\ R_{i}^{B} \end{pmatrix}$$

$$(1)$$

- İ: represents the internal nodes of the solid excluding the nodes situated on the slave boundary.
- C: represents the nodes belonging to the slave boundaries.
- **m**: represents the nodes belonging to the master boundaries.
- K^{α} : represents the matrix corresponding to the sub-domain Ω^{α} , $\alpha = A$; B.
- ΔU^{α} : represents the vector of the degrees of freedom of all the elements of Ω^{α} .

The continuity equation

The continuity equation at the master-master interface $\gamma_m^A-\gamma_m^B$ is written for the discrete case:

 $\{U_m^A\} - \{U_m^B\} = 0$

The equilibrium conditions

The equilibrium conditions at the master interfaces:

$$R_m^A\} - \{R_m^B\} = 0 \tag{3}$$

In a resolution context by sub-domains, the relation (??) will be satisfied by an iterative method. For this purpose, it is assumed that the resolution on each sub-domain was obtained by assuming a fixed value of $\{U_m^\alpha\}$ to estimate the displacement of the slave boundary using the system (??) for a resolution on each sub-domain:

$$\begin{bmatrix} K_{ii}^{\alpha} & K_{ie}^{\alpha} \\ K_{ei}^{\alpha} & K_{ee}^{\alpha} \end{bmatrix} \begin{pmatrix} U_{i}^{\alpha} \\ U_{e}^{\alpha} \end{pmatrix} + \begin{bmatrix} 0 \\ K_{em}^{C,\alpha} \end{bmatrix} \{ U_{m}^{\alpha} \} = \begin{pmatrix} R_{i}^{\alpha} \\ R_{e}^{\alpha} \end{pmatrix}$$
(4)

 $[K_{ee}^{\alpha}] = [K_{ee}^{S,\alpha}] + [K_{ee}^{C,\alpha}], \quad [K_{ie}^{\alpha}] = [K_{ie}^{S,\alpha}]$

 $[K^{Cea}_{ee}]$: the terms of the contact tangent matrix for the slave nodes. Knowing the value of $\{U^a_e\}$ for each sub-domain, it is possible to satisfy (3) via an iterative method, that is to obtain a correction to the solution to the $\{U^a_m\}$ field by:

$$\{R_m\} = \frac{1}{2} \sum_{\alpha=1}^{2} \{R_m^{\alpha}\}$$

$$\{U_m^{\alpha,i+1}\} \equiv \{U_m^{i+1}\} = \{U_m^i\} + \{\Delta U_m^i\}$$
$$\int \Delta U^i = \beta \sqrt{U^{A,i}} + (1-\beta) \sqrt{U^{B,i}}$$

$$\{\overline{U_{m}^{\alpha,i}}\} = -[K^{C,\alpha}]^{-1}\{R_{m}\}, \quad \alpha = A, B$$

$$\{R^{\alpha}_m\} = [K^{C,\alpha}_{m,e}]\{U^{\alpha}_e\} + [K^{C,\alpha}_{m,m}]\{U^{\alpha}_m\}$$

 $[K^{C,\alpha}]^{-1}$: correction matrix to be defined.

Definition of the correction matrix at the interface

The correction matrix at the interface can be defined in various ways. The following lines present two distinct techniques: Method 1

the direct method

Method 2

the method based the flexibility of the sub-domains.

APPLICATIONS: ONE-DIMENSIONAL THERMAL CONTACT

The length of the bars is L = 100 and its section A = 100. The thermal conductivity is K = 20.

Figure 3: Two bars in thermal contact, geometry of the contact interface

The topology of the global matrix taking into account the interface



Method 1: influence of the constant h



Method 2: influence of discretization

Number of nodes by subdomain	Number of iterations
4	25
10	66
15	101

CONCLUSIONS

Each sub-domain has a pair of slave and master boundariesThis approach should reduce the communication time between the processors during a parallel programming.

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Vers un Outil de Gestion Intégrée de l'Environnement: Impact des Changements Climatiques sur les Ressources en Eaux au Nord du Maroc (Cas du bassin versant d'Oued Boukhalef)

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Couplage de l'AC avec CC:

- Entrée des donnés physiques de terrain.
- Choix de climat de référence.
- Choix de scénarios climatiques.

Conclusion et perspectives

- Illustration des impacts de CC sur les ressources en eau
- Couplage avec les changements climatiques (CC)
- Modèle 2D du suivi du cycle de l'eau
- Envasement des barrages (localisation et nature)

• Entrée des données climatiques.

• Sortie des ressources en eau et traitement.

• Simulation du cycle de l'eau.

- Recharge des nappes (requiert une modélisation 3D)
- Hétérogénéité des sédiments transportés (modélisation 3D)

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Figure 9: Comparaison selon 4 scénarios